

37), but is often present conspicuously in hand-specimens (Fig. 36), and can be traced in increasing dimensions, till it connects itself with gigantic curvatures of the strata, which embrace whole mountains. These characters are sufficient to indicate a great difference between schistose rocks and ordinary stratified formations, in which the strata lie in continuous flat, parallel, and more or less easily separable layers.

In some instances, the folia can be seen to coincide with original bedding, as where a band of quartzite or of conglomerate is intercalated between sheets of phyllite or mica-schist. In such cases, there cannot be any doubt that the rock, though now more or less reconstructed and crystalline, was originally mere accumulated mechanical sediment. Many clay slates, phyllites, and mica-schists are obviously only altered marine clays, and some of them still retain their recognizable fossils. From such rocks, gradations can be followed into chistolite-schist, mica-schist, and fine gneiss. Quartzites and quartz-schists often still retain the false-bedding of the original sandy sediment of which they are composed. The pebbly and conglomeratic bands associated with some schists afford convincing proof of their original clastic nature. Thus, at the one end of the schistose series we find rocks in which an original sedimentary character remains unmistakable. At the other end, after many intermediate stages, we encounter thoroughly amorphous crystalline masses, that bear the closest resemblance to eruptive rocks into which they insensibly pass. In such instances, there can be little doubt that the amorphous structure is the original one, which has become schistose by subsequent deformation (Book IV. Part. VIII.). The banded arrangement of many coarse gneisses, however, may be an original segre-